Guest Lecture

Chris Landry, Executive Director Center for Snow and Avalanche Studies

Wednesday, April 18, 2:30 p.m. Building 1400 (Green Building), Conference Room 246

Topics:

- The Colorado Dust-on-Snow Program: an applied science effort that monitors the presence/absence of dust layers at mountain passes throughout Colorado. CODOS provides a series of "Update" analyses of how dust-on-snow is likely to influence snowmelt timing and rates during the snowmelt runoff season.
- Other Center Activities

The **Center for Snow and Avalanche Studies** is an independent Colorado not-for-profit corporation founded in 2003 and located in Silverton, CO. The Center serves the mountain science community and regional resource managers by hosting & conducting interdisciplinary research and monitoring that captures weather, snowpack, radiation, soils, plant community, and hydrologic signals of regional climate trends. Programs include Mountain System Monitoring, Colorado Dust-on-snow, Interdisciplinary Research, and Field and Education Workshops.

Chris Landry began skiing at the age four at Big Mountain, Montana. His skiing career progressed from fledgling jumps to a junior and college racing career. Ski racing eventually turned into a passion for climbing, ski mountaineering, and even a few pioneering ski descents in North America. Along the way, Chris pursued instruction in avalanche safety, became an instructor for the American Avalanche Institute, published a field workbook for backcountry skiers and guides, and established a practice as a private avalanche consultant/forecaster. Most recently, Chris earned a MS in the Department of Earth Sciences at Montana State University -Bozeman, where he researched the spatial variability of snow stability on uniform slopes.



Research and Monitoring

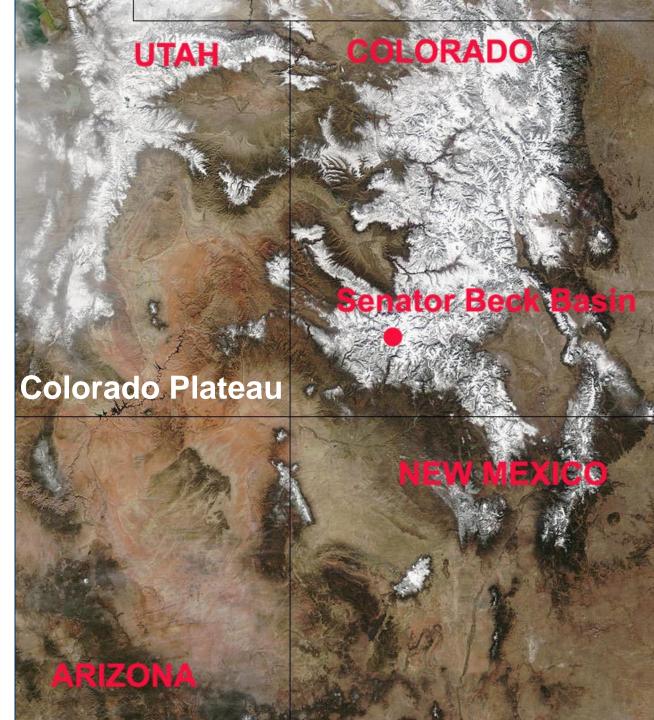
Mountain System Processes and Change

Senator Beck Basin Study Area

Chris Landry

Center for Snow and Avalanche Studies Silverton, CO

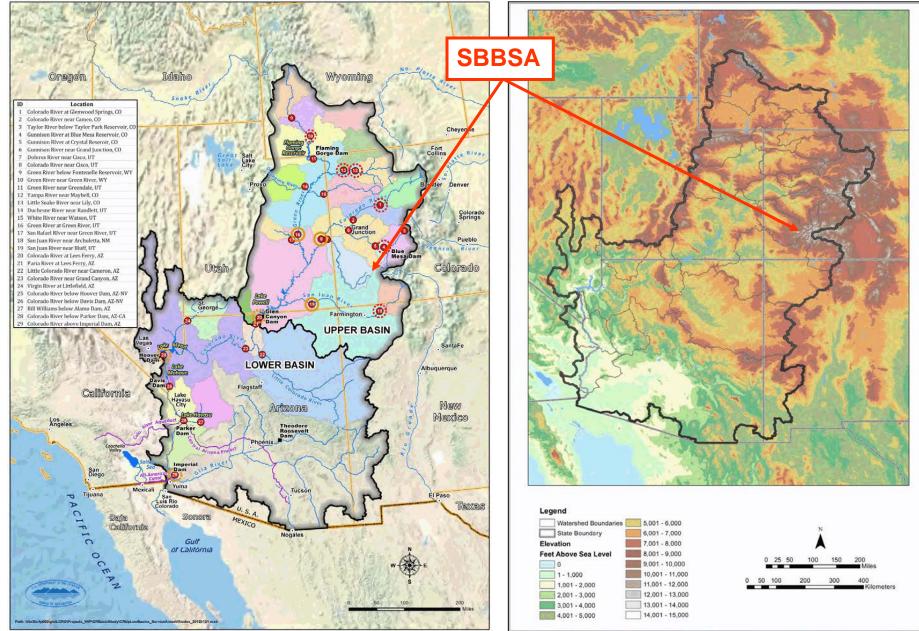


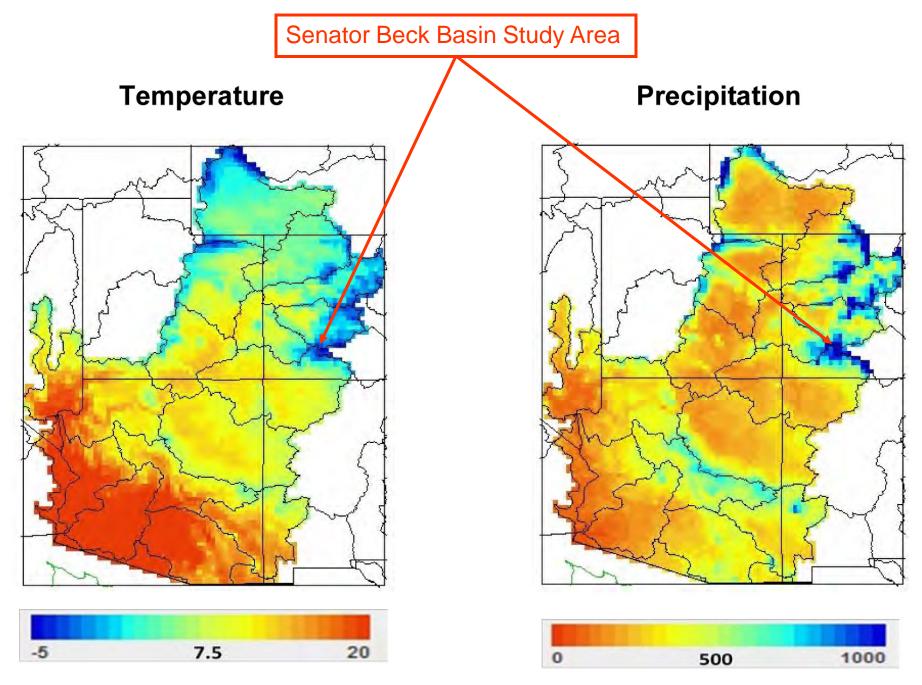


Senator Beck Basin Study Area – Red Mountain Pass from Putney Study Plot – 12,325'

Senator Beck Basin Study Area – Sentry Site for Upper CRB

CRB Water Supply & Demand Study: Tech Report B - pg's B7 & B13

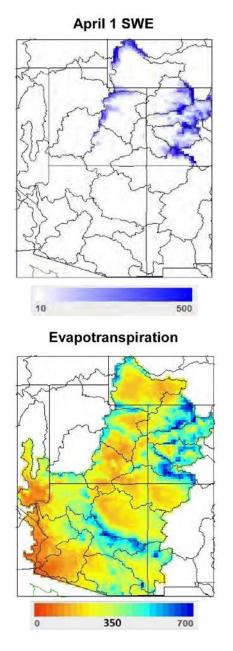


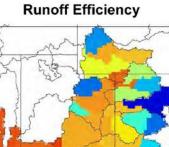


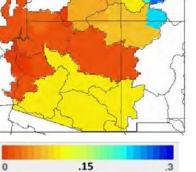
CRB Water Supply & Demand Study: Tech Report B - pg B14

FIGURE B-8

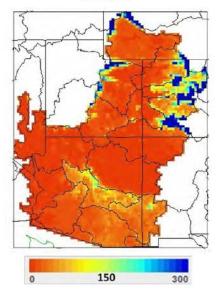
Estimated Average Annual ET and Runoff (millimeters), April 1 SWE (millimeters), and Annual Average Runoff Efficiency (fraction of precipitation converted into runoff) for 1971–2000 Derived from historical VIC simulations.







Runoff

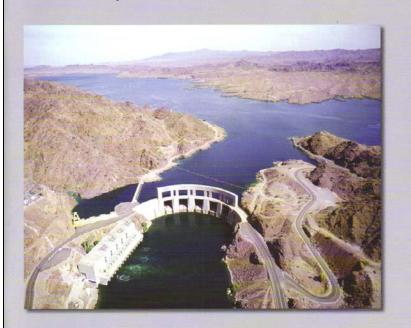


CRB Water Supply & Demand Study: Tech Report B - pg B18



Climate Change and Water Resources Management: A Federal Perspective

US Army Corps



Circular 1331

U.S. Department of the Interior U.S. Geological Survey

2 Tracking Climate Change Impacts

Monitoring data is essential for understanding and "Wishing the impacts of climate change. This chapter seeks to uniters the following questions:

- How are monitoring data used to track climate impacts?
- · How do data inform physical system understanding?
- · What monitoring networks currently exist?

2.1 Tracking Hydrologic Change: Monitoring Networks

Current projections of climate changes and their potential intracts harbor many uncertainties, and these uncertainties are enticled to dissipate in the near term. Within these uncertainties are the possibility for surprises, which could be unpleasant and guids to appear. In this context, a strategy that balances detecting and adjusting to changes against extrapolating (including modeling) and anticipating changes will be most prudent. Thus, monitoring of elimatic and hydrologic conditions plays an important role in addressing potential climate changes.

To detect hydrologic changes due to climate change or other causes, data from long-term monitoring networks are sential for establishing baseline conditions and tracking any plunges over time. Monitoring networks are also essential for fully understanding the hydrologic processes that lead to charges in water resources and for calibrating and validating works used to project future conditions. In turn, information doon possible or likely future changes to climate improves the event news of planning studies and allows the development and implementation of reasonable strategies for adapting to a changing climate.

Key Point 4: Long-term monitoring networks are critical for detecting and quantifying climate change and its impacts. Continued improvement in the understanding of climate change, its impacts, and the effectiveness of adaptation or minigation actions requires confinited operation of existing long-term monitoring networks and improved senses deployed in space, in the antrosphere, in the occurs, and on the Earli's surface.

Monitoring networks include in situ methods as well as remote sensing technologies such as radar and satellites. Existing data allow us to look at data retrospectively. However, unonitoring networks must continue to operate into the future if we ure to detect future changes in hydrologie systems due to elimate change (or the lack thereof) and to eraft effective responses.

To be useful for climate change studies, monitoring networks need to be in place in locations relevant to water managers. For example, monitoring stations should be located in watersheds important for water supply or vulnerable to

changes in water quality. In addition to monitoring of the natural system, data on human water use can be valuable in planning for climate change. The USGS periodically publishes estimates of water use in the United States by sector (*dar* example, Ilusion and others, 2001) compiled (rom data collected by State and local agencies. The periodic nature of these reports and the varying data-collection methods limit their unity for evaluating demand interactions with climate.

Climate change is easier to detect on global to regional scales. Monitoring networks for detacting change are especially valuable when they are regional or involve local networks that are integrated to allow regional analyses. Also needed for planning and operational analysis is a comprehensive set of parameters that characterize current and forme climate conditions.

A number of Federal. State, and local agencies operate observation networks that are valuable for climate change analysis. The USGS operates the largest water monitoring network in the United States, as well as biological-monitoring networks. These are briefly described in the inset box. NOAoperates the Nation's largest metoorological network and provides data on oceans. The NOAA observational networks are also described in an inset box. Other Federal agencies also maintain important water-monitoring networks, such as the Natiral Resources Conservation Service's snow surveys and Snowich network. State and local agencies are able to supplement these larger networks with needed local data. USACE and Reclamation also conduct project-specific water resources-monitoring activities.

Key Point 5. Monitoring needs to focus on locations that describe the climate signal the example, upstream and downstream from major water-management mirastructure or in vulnerable ecological reaches).

2.2 Tracking Hydrologic Change: Trend Analysis

As discussed in chapter 1, climate change is expected to cause changes to streamflow, precipitation, and other hydroclimatic variables. The continuous long-term streamflow and meteorological records described in the preceding zection are critical for detecting trends or chifts in the statistics, of historical streamflow or other hydroclimatic variables. Such nonstationarity in hydroclimatic conditions would represent a change from the assumptions that have been used to design and manage water resource systems. Consequently, it is important to know if and how trends manifest themselves.

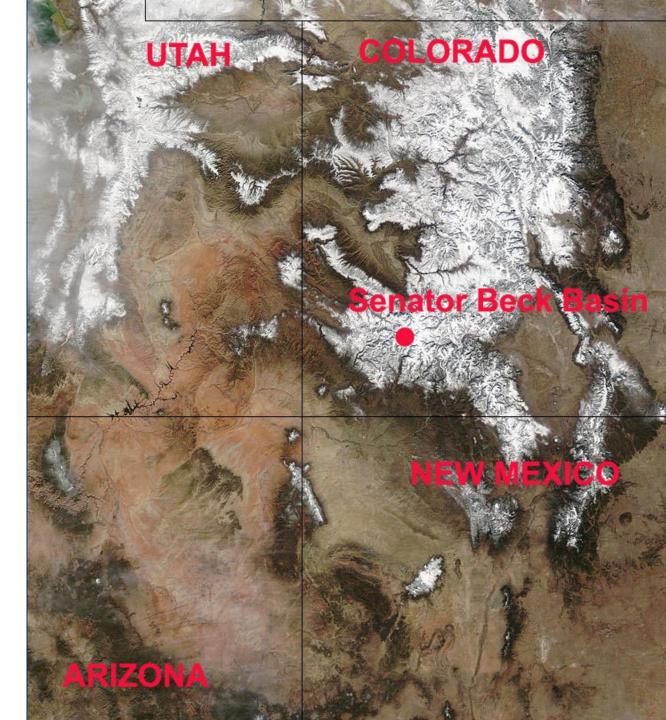
Trend detection must be carried out with care, as trends may also be caused by land use changes, changes in water infrastructure, or other factors. Furthermore, which the magnitude of a trend may be relatively cary to quantify, its statistcal significance may be more ambiguous because of natural climate variability and long-term persistence, which can cause oscillatory patterns in long-term hydroclimatic records (Cohn and Lins, 2005).

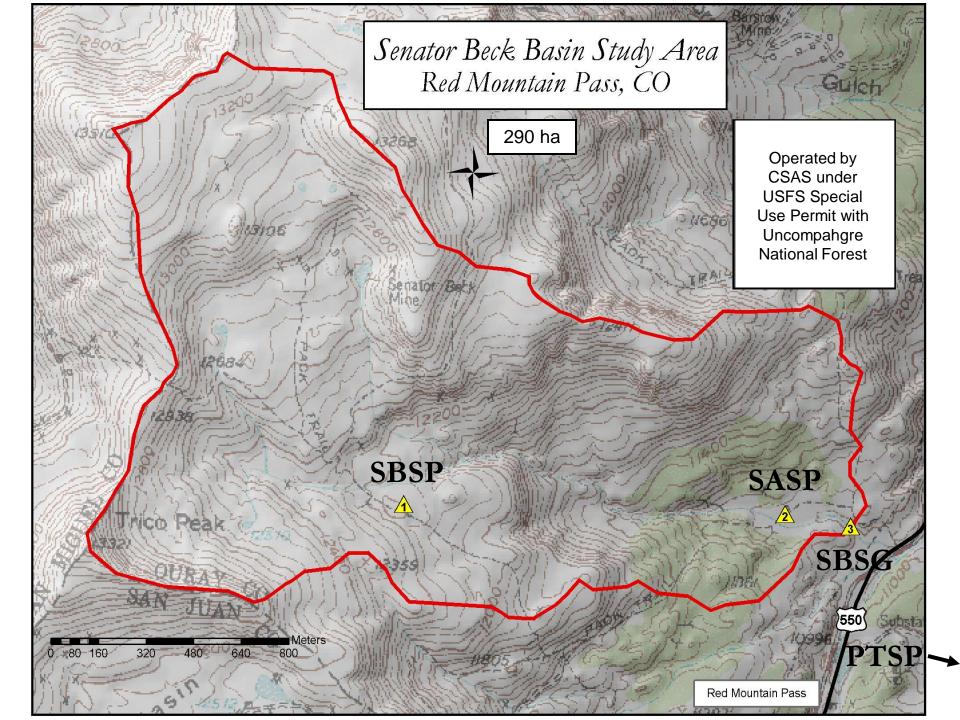
2 Tracking Climate Change Impacts 13

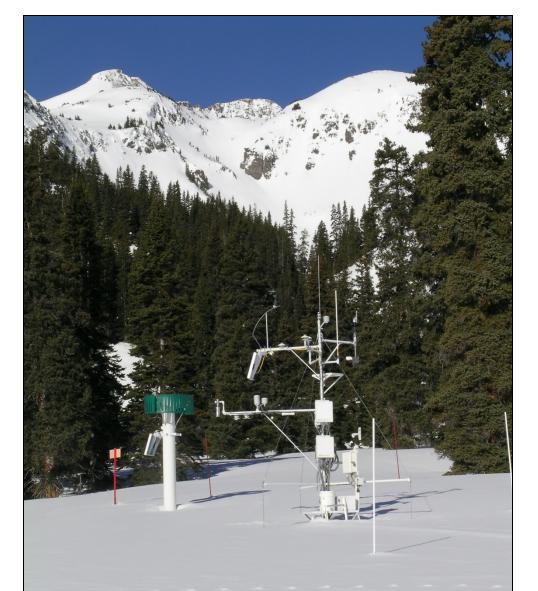
SBBSA

Sentry Site for change in Upper CRB ...

- Snowmelt Runoff
- Mountain Precip
- Mountain Temps
- Mountain Winds
- Mountain Radiation
- Mountain Vegetation
- Mountain Soil



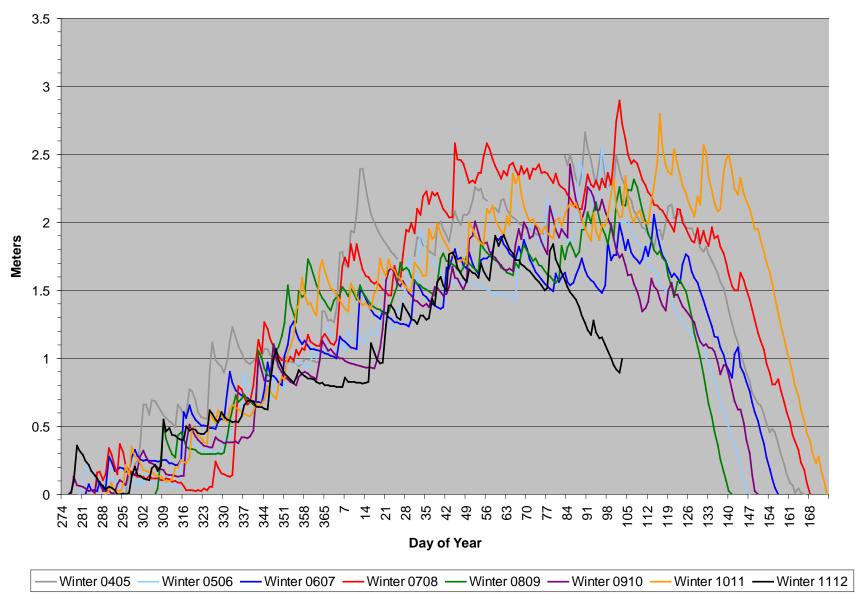


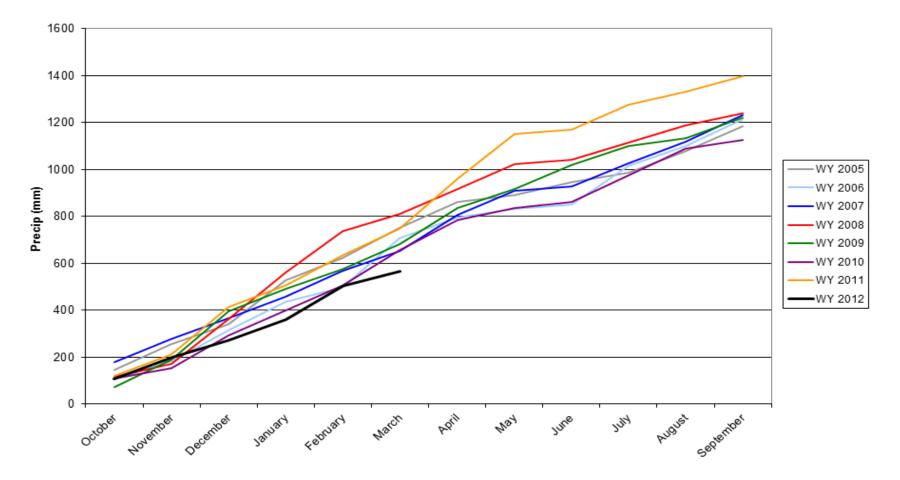


Swamp Angel Study Plot 11,050' (3368 m)

SASP Instrumentation 6 m Mast CR10X Dataloggers (2), Multiplexer (1) ETI Precipitation Gauge Wind Speed & Direction (2) Air Temp and RH (2) Barometric Pressure Height of Snow Broadband SW (2 up, 1 down, shadow array) NIR SW (1 up, 1 down) Pyrgeometer (1 up) Infrared Snow Surface Temp Snow Temperature (5) Soil Temperature (4) Soil Volumetric Water Content Soil Heat Flux

Height of Snow - Swamp Angel Study Plot as of 2400 hours





Water Year Cumulative Precipitation at End of Month Swamp Angel Study Plot - Senator Beck Basin Study Area at Red Mountain Pass

Annual Precipitation Quantity, Timing, and Phase Senator Beck Basin Study Area - Swamp Angel Study Plot

Water Year Monthly Means	Month	Mean Winter Storms	Mean Days Precip	Mean Total Precip - mm	Mean YTD Precip - mm	Mean Snow precip - mm	Mean Rain precip - mm
	October	2.2	11.4	108.3	108.3	79.6	28.8
Period of Record	November	3.3	11.6	93.1	201.4	90.9	2.2
= WY 2004 - WY 2012 (9 yrs)	December	3.7	16.1	137.0	338.4	135.2	1.8
= WY 2004 - WY 2011 (8 yrs)	January	3.4	13.1	117.2	455.7	117.2	0.0
= WY 2005 - WY 2011 (7 yrs)	February	3.3	15.9	118.2	573.9	118.2	0.0
	March	3.1	14.4	109.7	683.6	109.7	0.0
	April	4.4	15.0	138.4	836.6	138.4	0.0
	May	2.4	10.9	77.0	913.6	68.5	8.5
	June	0.4	6.6	36.0	973.1	6.1	29.9
	July	0.0	14.6	97.3	1070.4	0.0	97.3
	August	0.0	15.0	77.9	1148.3	0.0	77.9
	September	0.1	12.1	82.1	1230.4	12.7	69.4

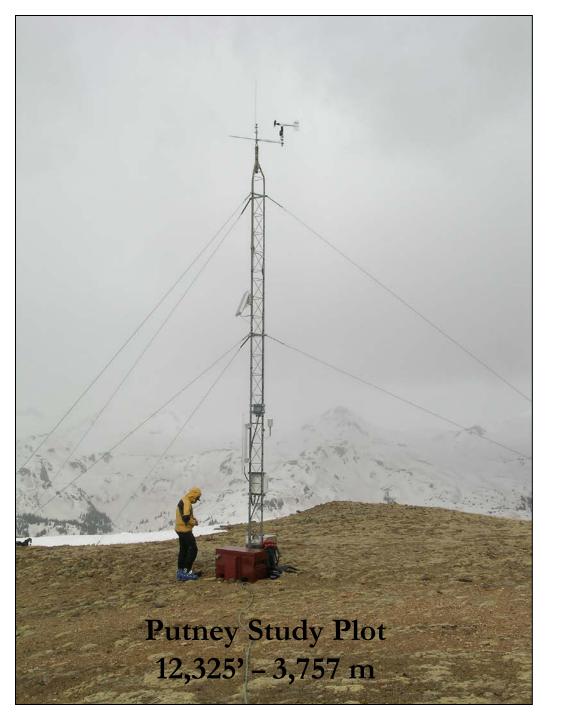
Water Year Annual Means	Winter Storms	Days Precip	Total Precip - mm	YTD Precip - mm	Snow precip - mm	Rain precip - mm
Annual Means	26.3	161	1,230	1,230	910	321
Period of Record - Water Years	2004-2011	2005-2011	2005-2011	2005-2011	2005-2011	2005-2011
					73.9%	26.1%



SBSP Instrumentation

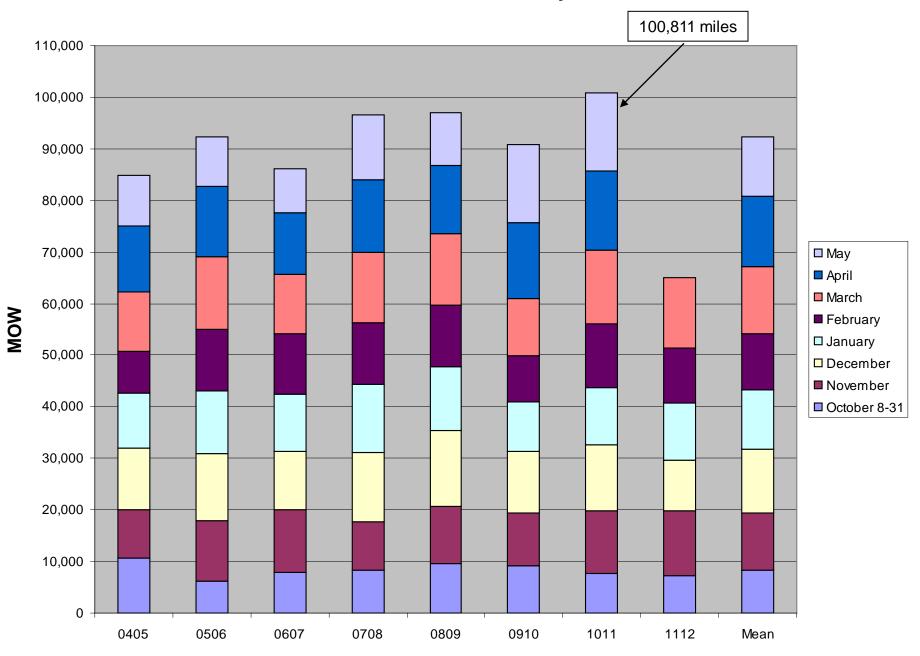
10 m Mast Campbell CR10X Dataloggers (2), Multiplexer (1) Wind Speed & Direction (2) Air Temp and RH (2) Height of Snow Broadband SW (2 up, 1 down, shadow array) NIR SW (1 up, 1 down) Pyrgeometer (1 up) Infrared Snow Surface Temp Snow Temperature (5) Soil Temperature (4) Soil Volumetric Water Content Soil Heat Flux

Snow Profile Plot

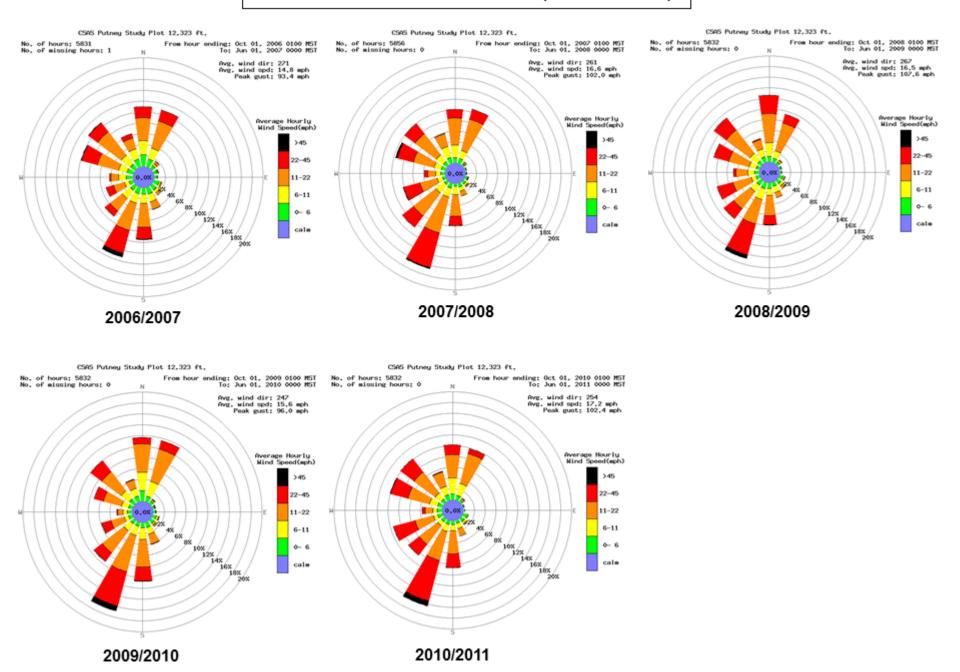


PTSP Instrumentation

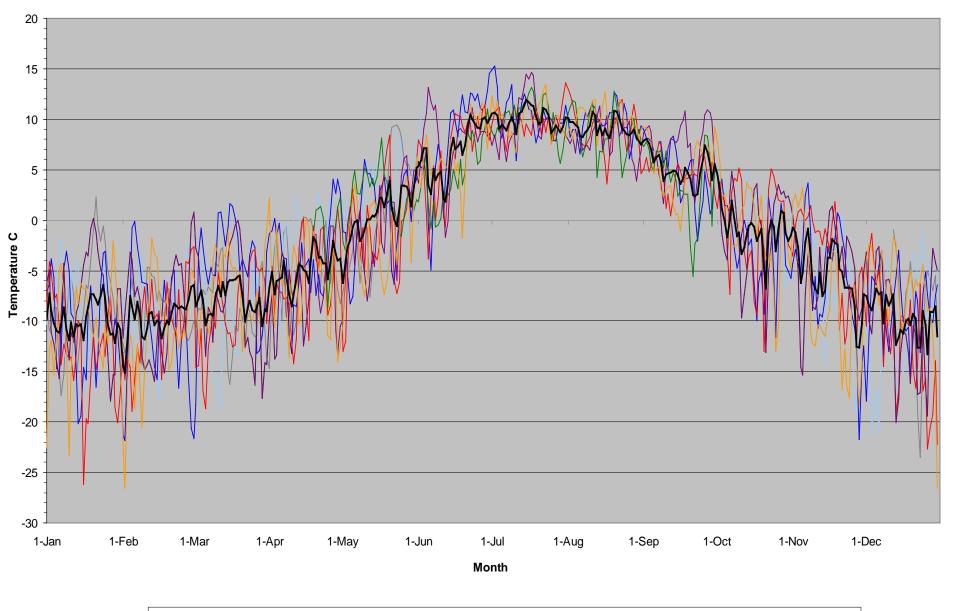
10 m Mast Campbell CR10X Datalogger Wind Speed & Direction Air Temp and RH Total Miles of Wind at PTSP by Season



PTSP Winter Wind Roses (10/1 – 5/31)



Putney Study Plot 24-Hour Mean Air Temperatures Elevation 12,325'

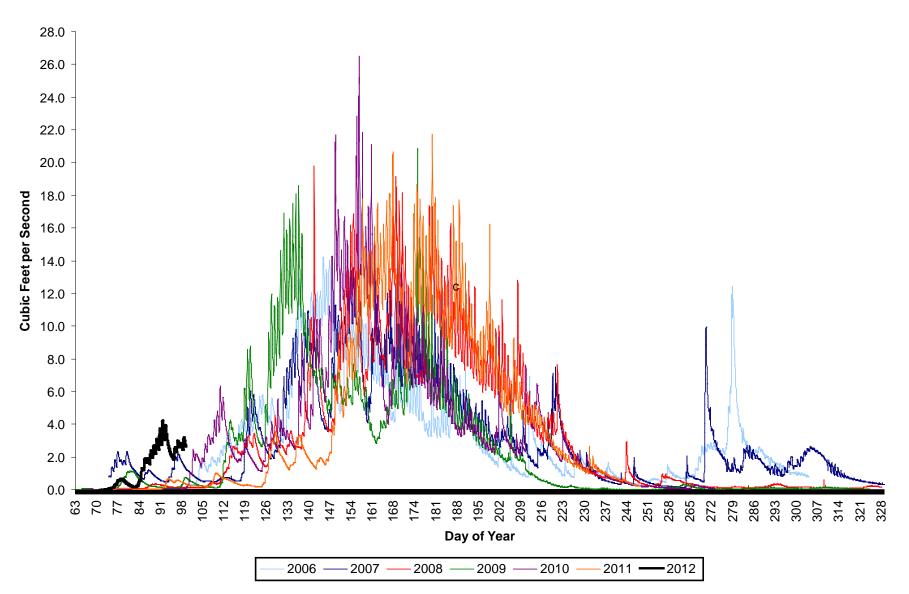


SBSG Instrumentation

Broad-crested, notched weir 0.1 – 30 cfs capacity Campbell CR10 Datalogger Stage – Druck transducer Stage – staff gauge Water Temp and Conductivity

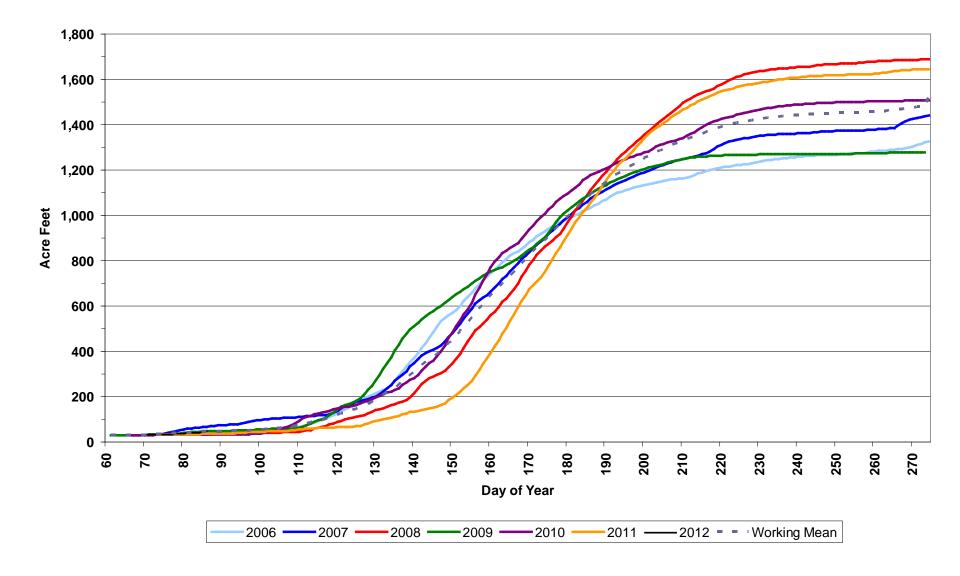
> Senator Beck Stream Gauge 11,030' – 3,362 m

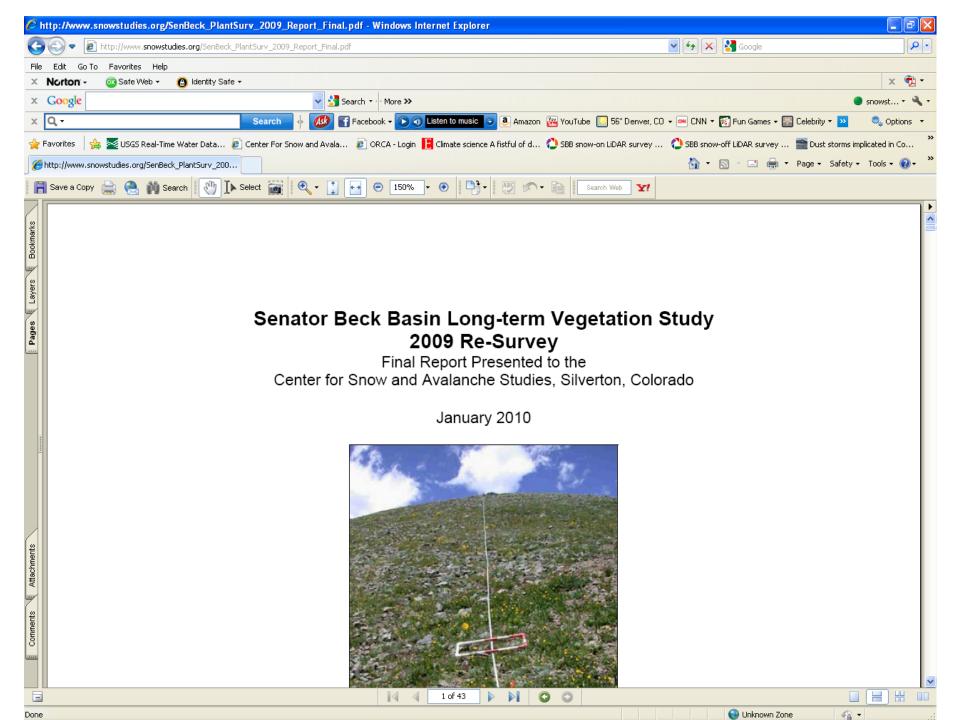
Senator Beck Basin Hourly Discharge WY 2006, 2007, 2008 & 2009, 2010, 2011, 2012



Senator Beck Basin Cumulative Discharge - 2006 to 2012

as measured at Senator Beck Stream Gauge (SBSG)





Mountain System Monitoring

Monitoring the plant community as a bellwether for regional climate 'state' in 5-year repeat studies

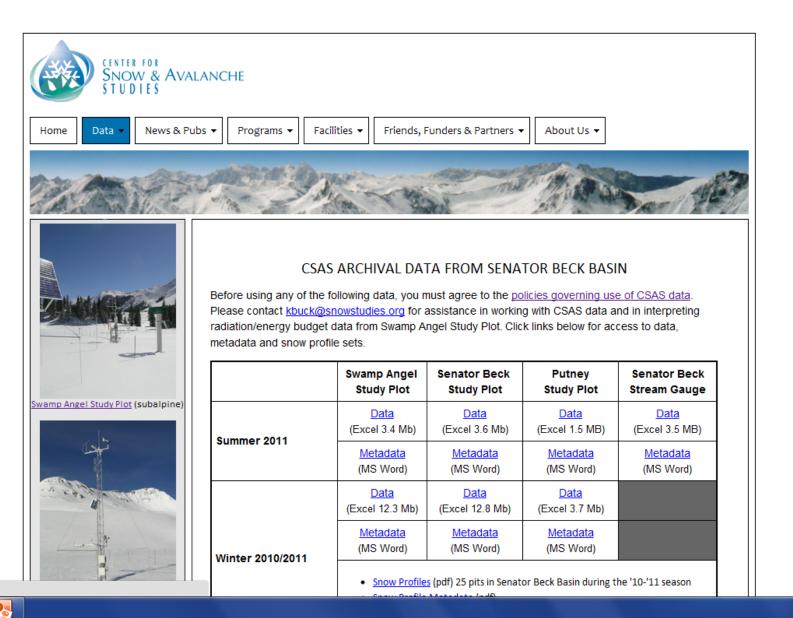
Vegetation Change =

- Change in Snowcover,
- Change in ET,
- Change in Albedo,
- Change in Runoff



August 2009





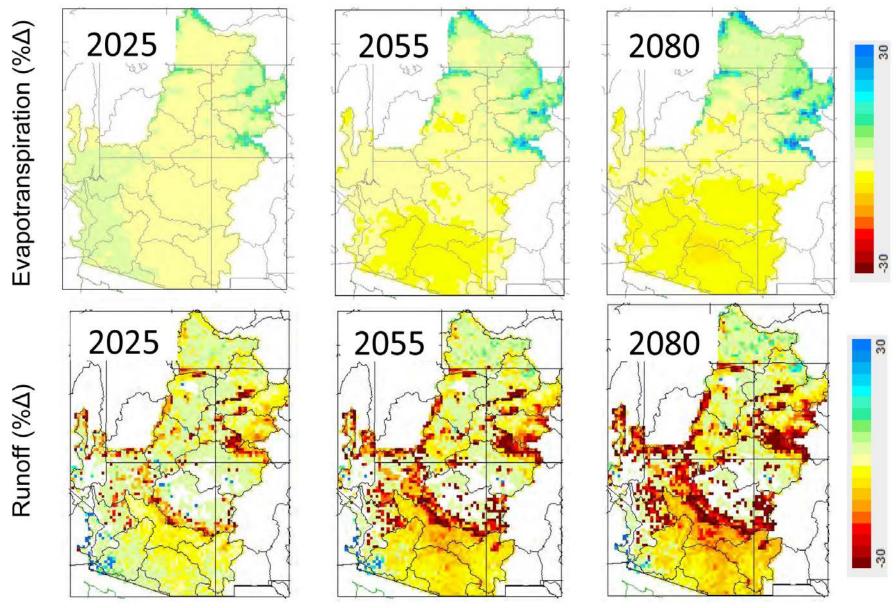
bsp1.html



FIGURE B-40

Mean Projected Percent Change in Annual ET and Median Projected Percent Change in Runoff⁸

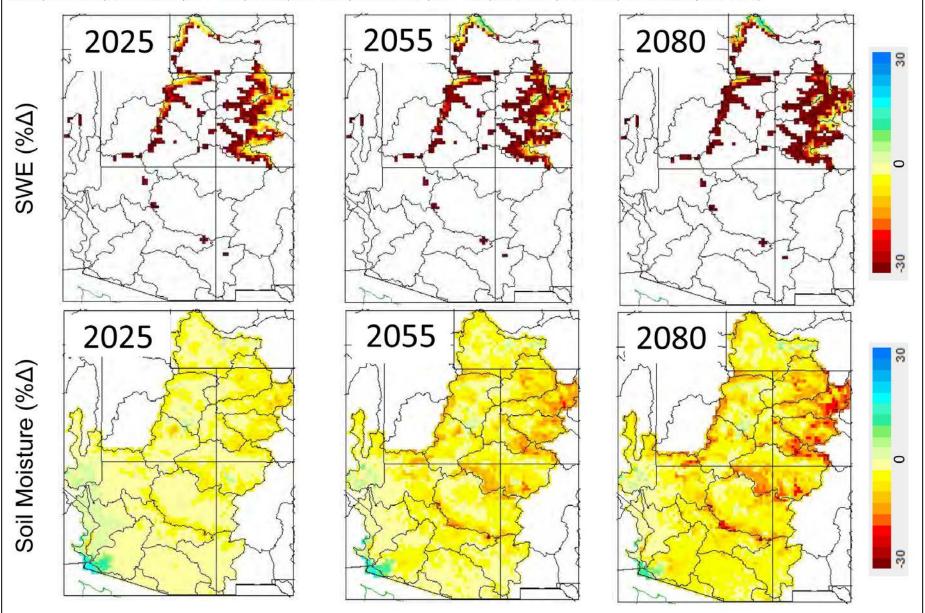
2025 (2011 – 2040) versus 1985 (1971-2000), 2055 (2041 – 2070) versus 1985 (1971-2000), and 2080 (2066 – 2095) versus 1985 (1971-2000).



CRB Water Supply & Demand Study: Tech Report B - pg B58

FIGURE B-41

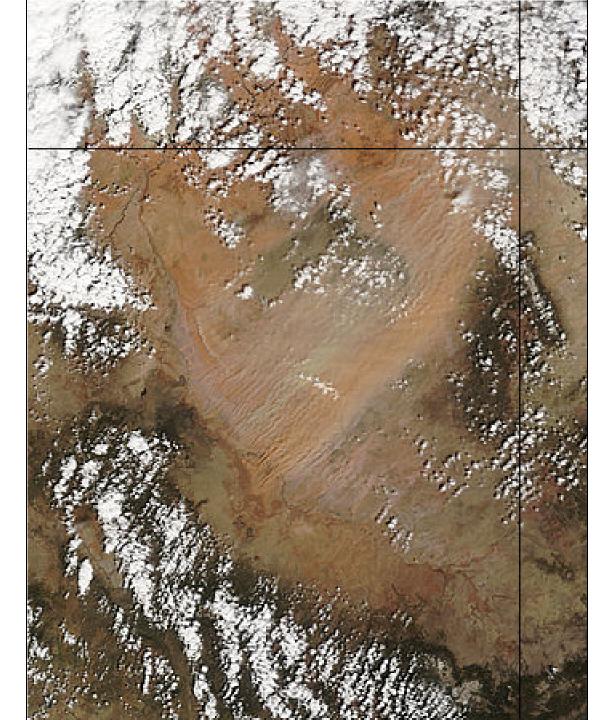
Mean Projected Percent Change in April 1 SWE and July 1 Soil Moisture 2025 (2011–2040) versus 1985 (1971–2000); 2055 (2041–2070) versus 1985 (1971–2000); and 2080 (2066–2095) versus 1985 (1971–2000).



CRB Water Supply & Demand Study: Tech Report B - pg B59

Senator Beck Basin Study Area, CSAS, and dust-on-snow science

May 12, 2009 – from Peak 13,510' at top of Senator Beck Basin Study Area looking southwest



D8 – WY2009 Little Colorado River April 3, 2009



D8 – WY2009 – Senator Beck Basin

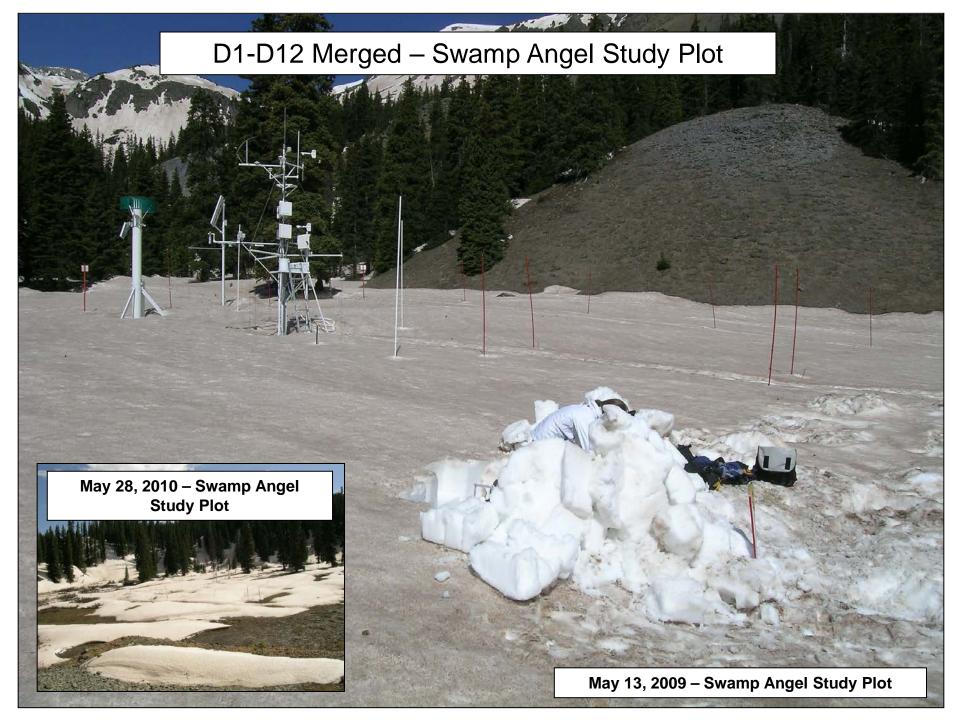
CSAS Putney Site 12,327ft. No. of hours: 18 From: Apr 03, 2009 1200 MST To: Apr 04, 2009 0600 MST Ν Avg. wind dir: 203 Avg. wind spd: 29.9 mph Peak gust: 72.9 mph Wind Speed (mph) >25 15-25 0.0% Е 5-15 Ы 5% 15% 20% 25% 30% 10% 0-5 calm 35% 40% 45% 50% S

Senator Beck Basin: March and April 2009 Dust Layers



SASP – April 22, 2009

SBSP – April 24, 2009

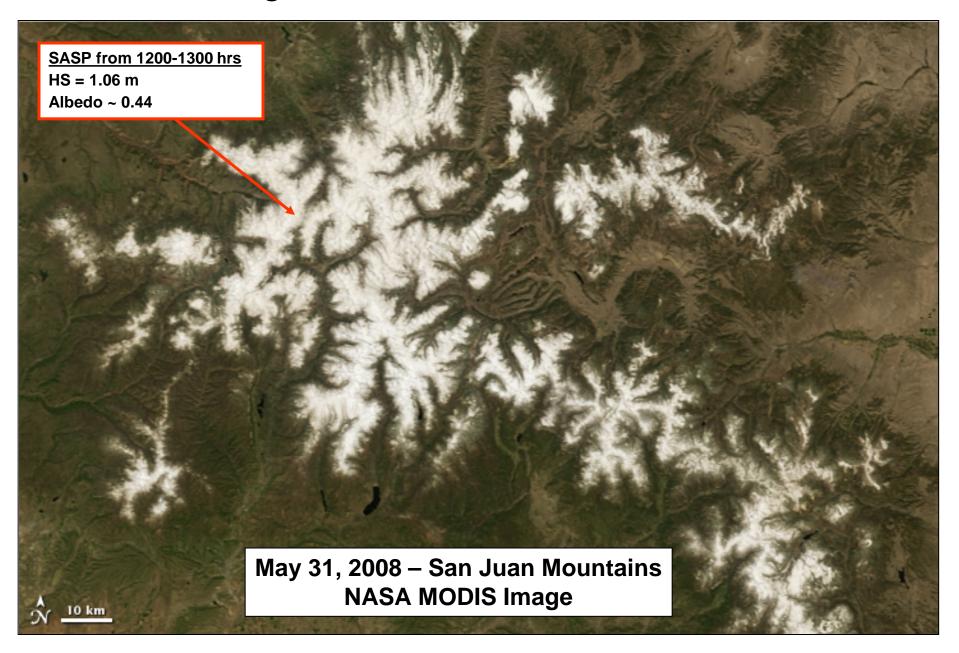


Storm #34 – June 20, 2011

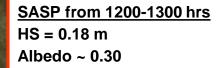
Dust Layers #11-4, merged

Snow Profile #25 - June 21, 2011 Senator Beck Study Plot

Large-Scale Albedo Reductions



Large-Scale Albedo Reductions



May 18, 2009 – San Juan Mountains NASA MODIS Image

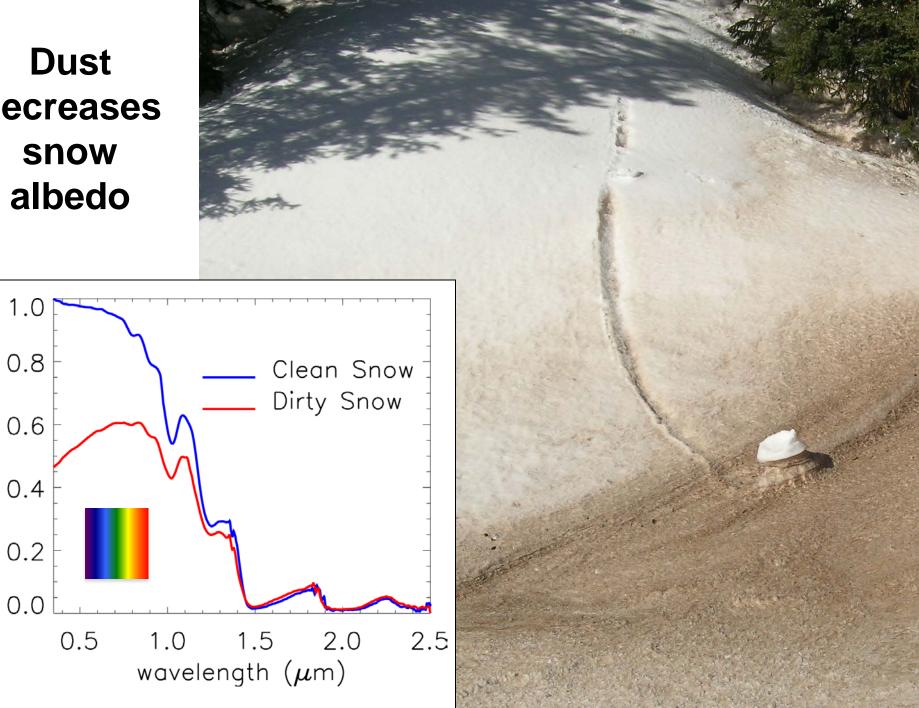
Dust-on-Snow Events Documented per Month, by Winter

Senator Beck Basin Study Area at Red Mountain Pass – San Juan Mountains

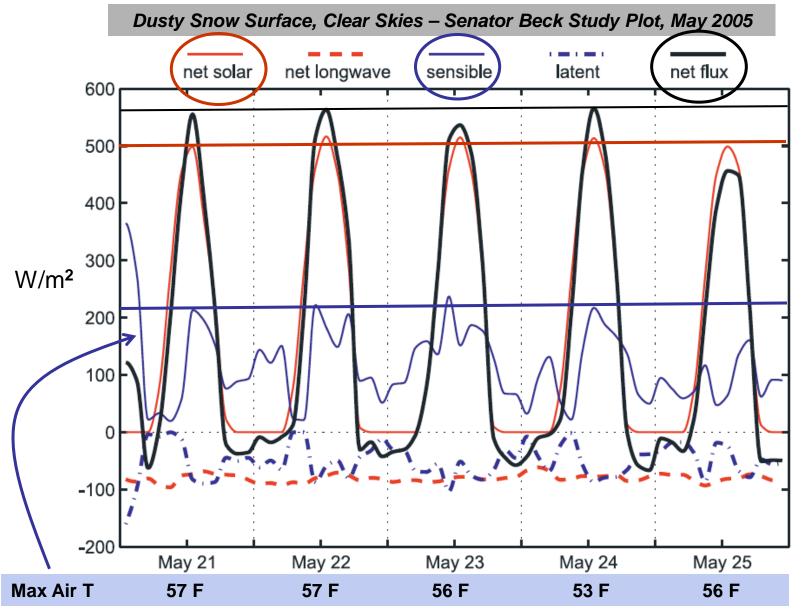
	Oct	Nov	Dec	Jan	Feb	Mar	Apr	Мау	Jun	Total
2002/2003					2		1			3
2003/2004							2	1		3
2004/2005	0	0	0	0	0	1	2	1	0	4
2005/2006	0	0	1	0	1	1	3	2	0	8
2006/2007	0	0	1	0	1	1	3	1	1	8
2007/2008	0	0	0	0	0	3	3	1	0	7
2008/2009	1	0	1	0	1	4	5	0	0	12
2009/2010	1	0	0	0	0	1	4	3	0	9
2010/2011	0	0	0	0	1	3	3	4	0	11
2011/2012	0	2	1	0	0	3	2			8 to-date

Dust decreases snow albedo

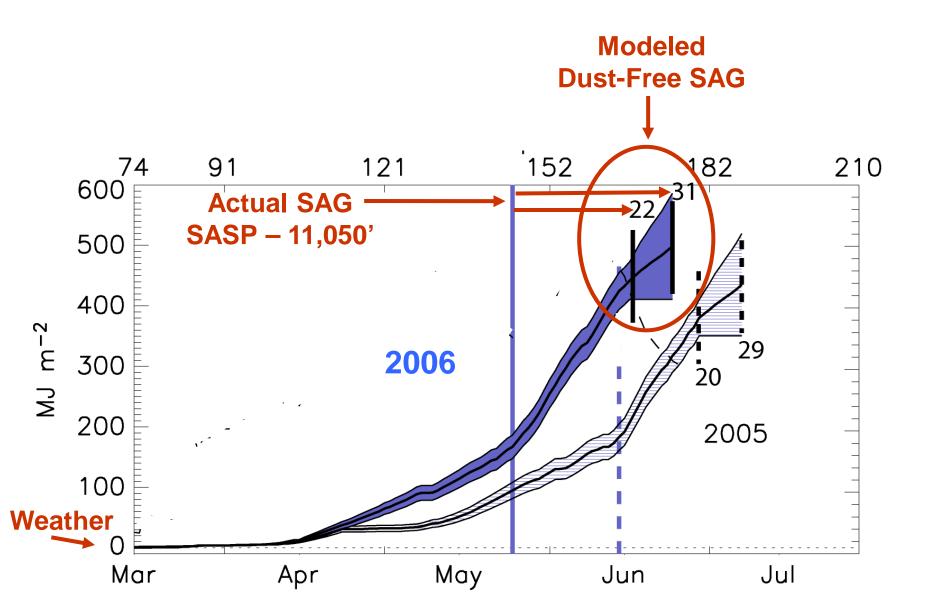
HDRF

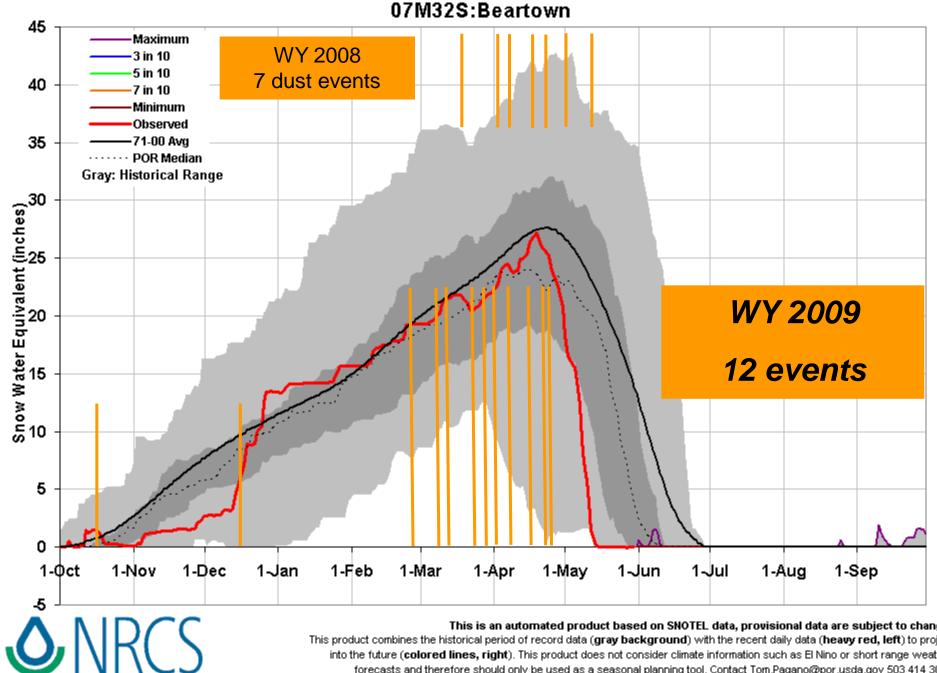


Enhanced Snowmelt Energy Input

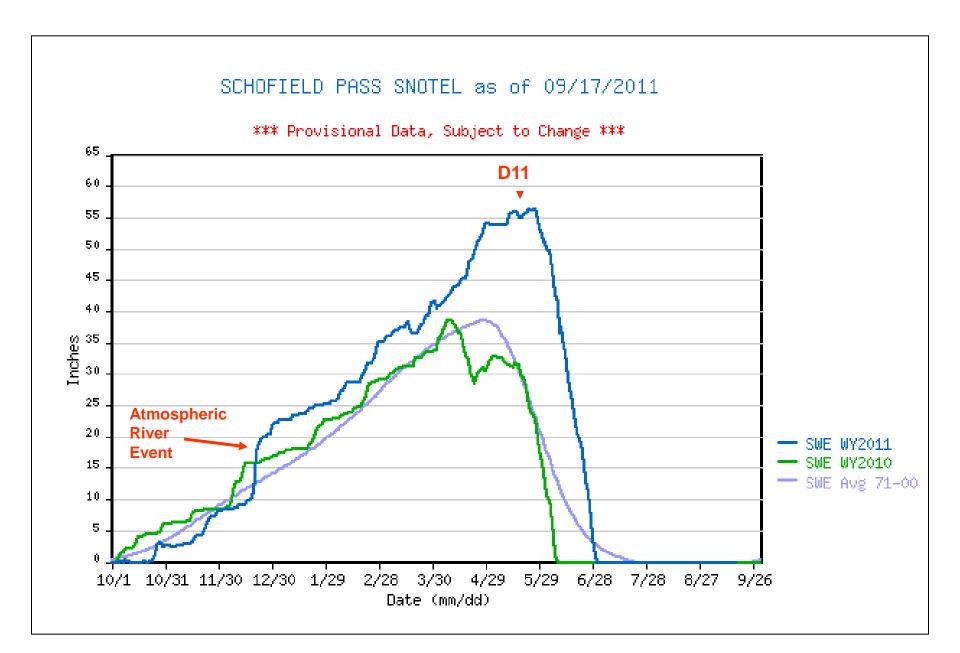


Radiative Forcing by Dust



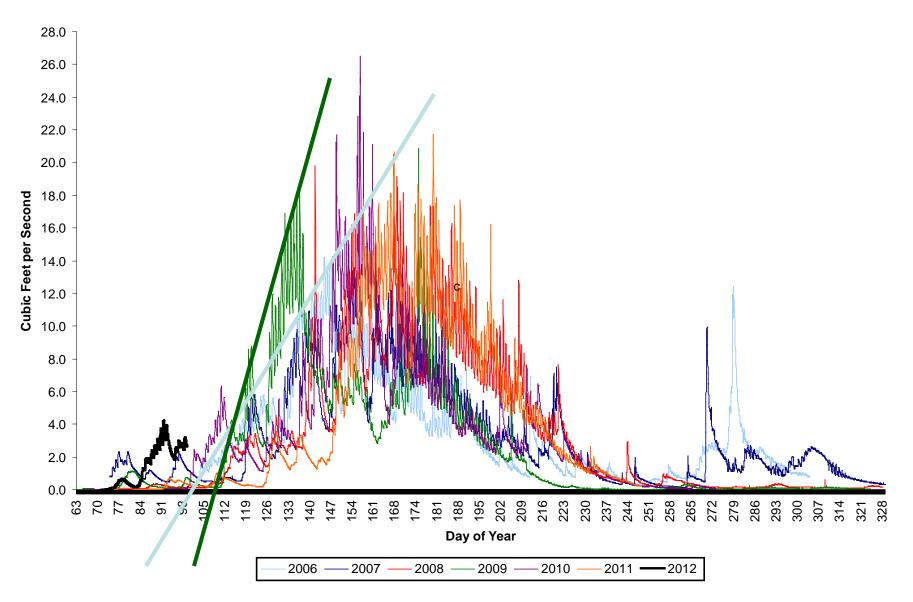


This is an automated product based on SNOTEL data, provisional data are subject to change. This product combines the historical period of record data (gray background) with the recent daily data (heavy red, left) to project into the future (colored lines, right). This product does not consider climate information such as El Nino or short range weather forecasts and therefore should only be used as a seasonal planning tool. Contact Tom.Pagano@por.usda.gov 503 414 3010



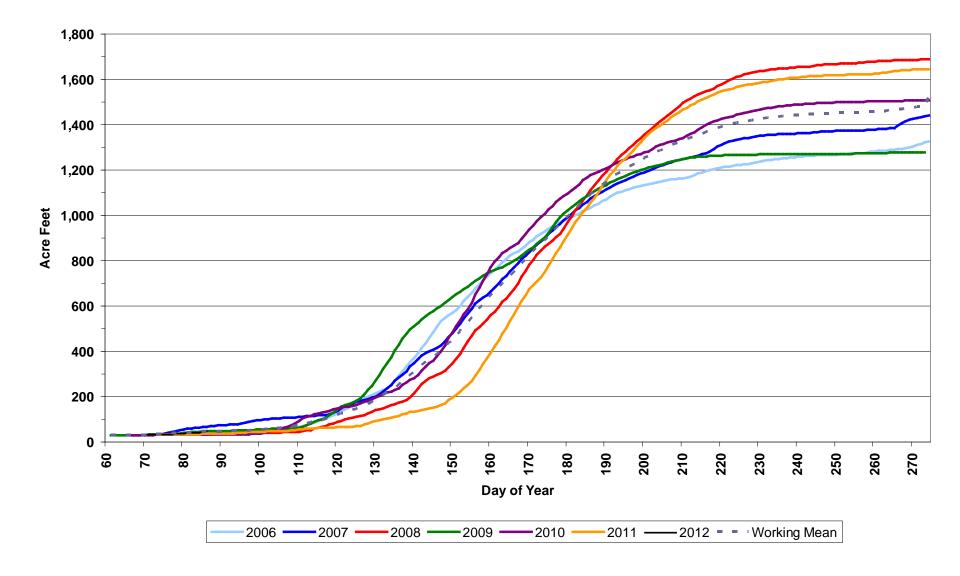


Senator Beck Basin Hourly Discharge WY 2006, 2007, 2008 & 2009, 2010, 2011, 2012

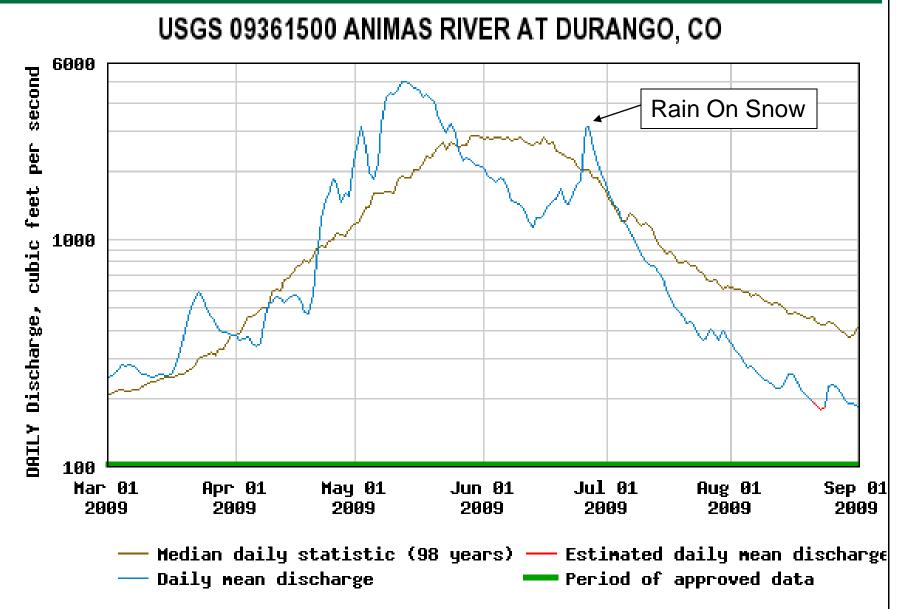


Senator Beck Basin Cumulative Discharge - 2006 to 2012

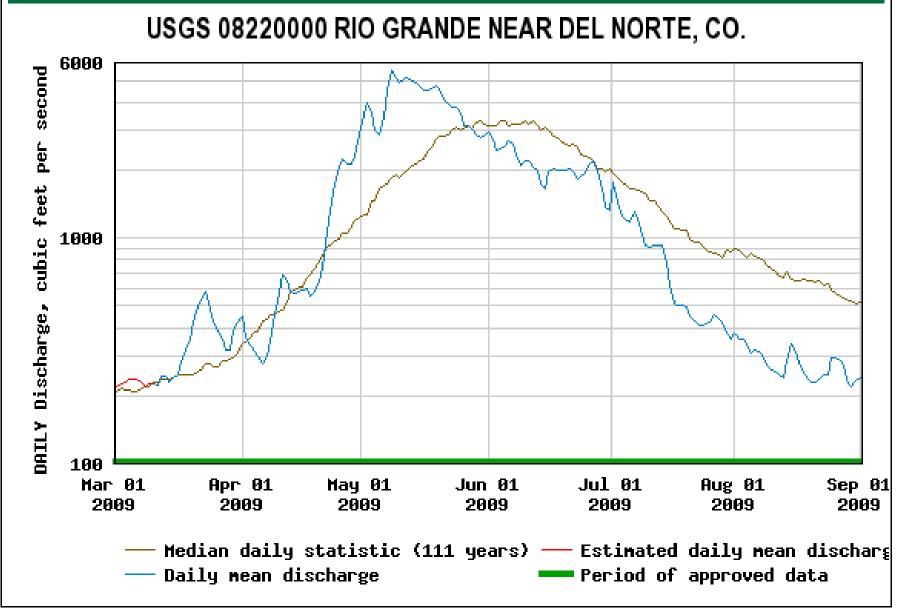
as measured at Senator Beck Stream Gauge (SBSG)



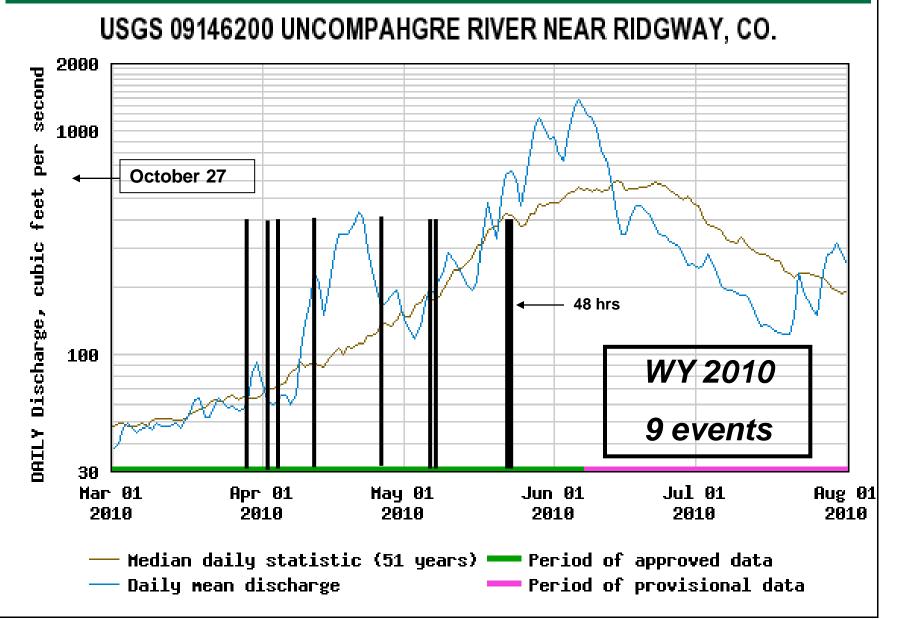
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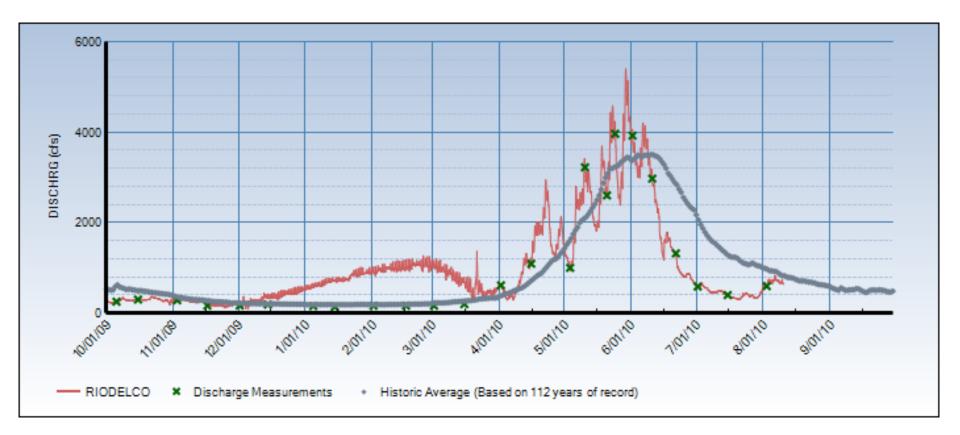
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≊USGS

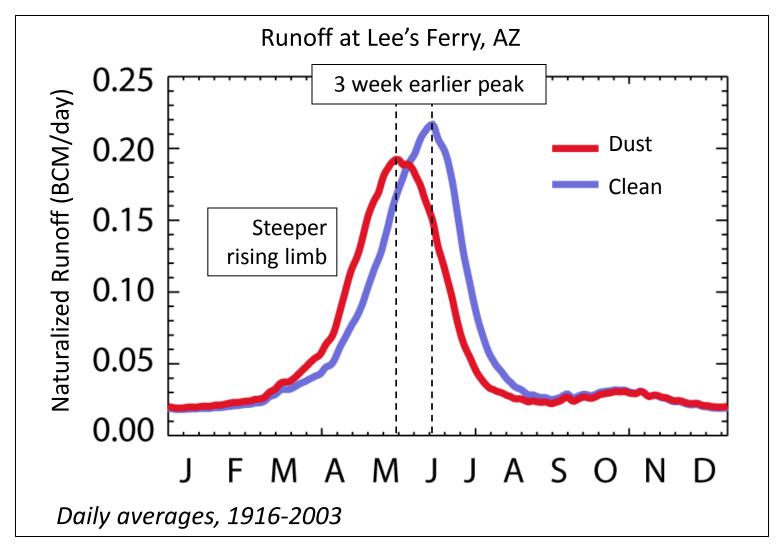


Rio Grande at Del Norte WY 2010



Dust-on-Snow Shifts Upper CRB Hydrograph*

*not including 2009, 2010 dust seasons

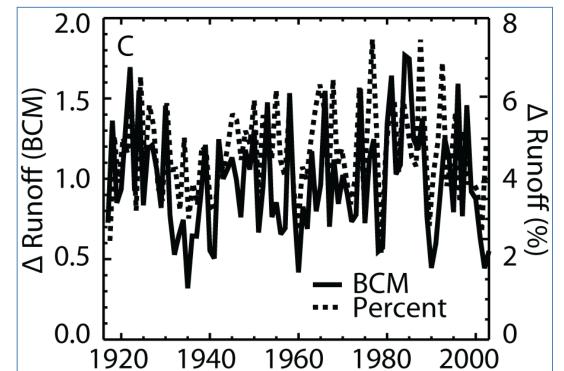


Painter, Deems, et al., PNAS (2010)

At the scale of the Upper CRB, modeling shows:

DOS = Earlier SAG = Increased ET = Reduced Runoff

- Mean ∆ Runoff: -4.9%
 -811,000 acre-ft
- Range:
 -2.3 to -7.6%
 -243k to -1,460k AF



*based on pre-2009 dust loading

Painter, Deems, et al., PNAS (2010)

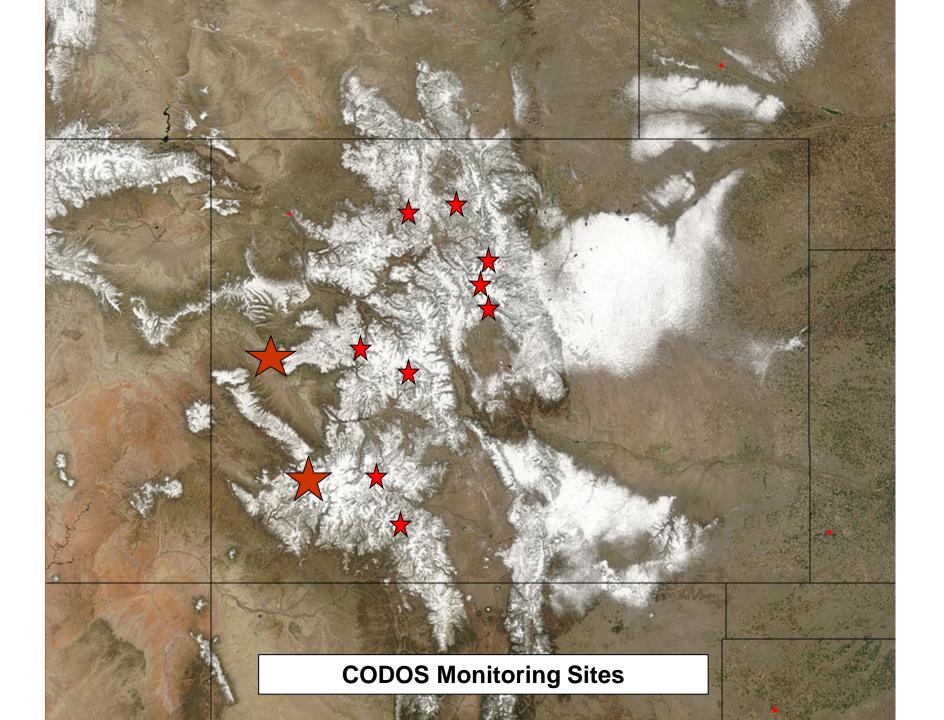
CSAS Colorado Dust-on-Snow (CODOS) Program

Timely, iterative monitoring and analysis of dust effects on snowmelt timing and rates throughout the Colorado mountains ... supplemental to CBRFC forecasts.

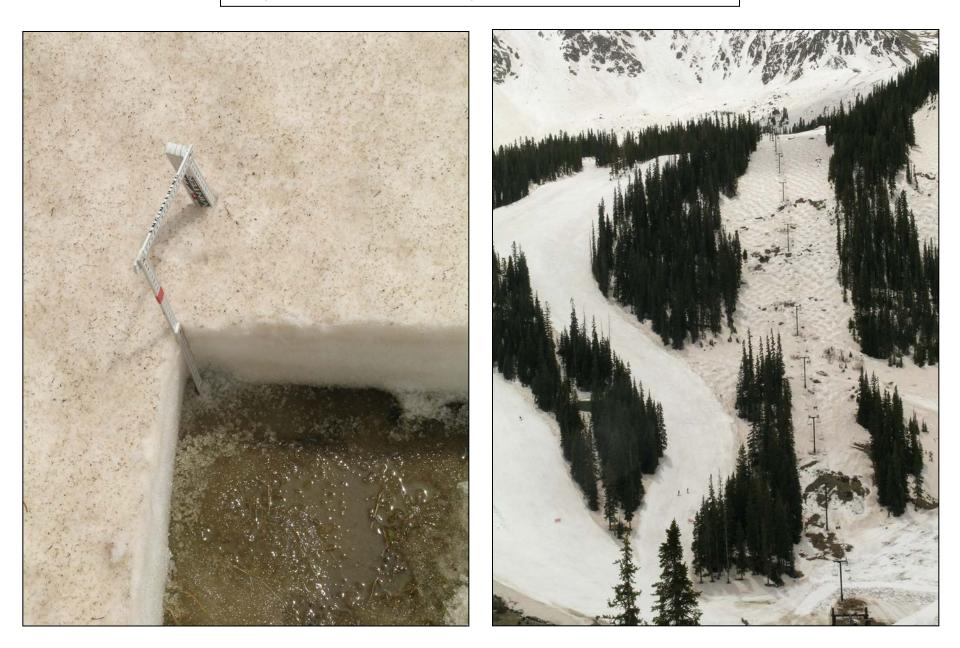
Spring 2009 – From Senator Beck Basin Study Area

CSAS Colorado Dust-on-Snow Program

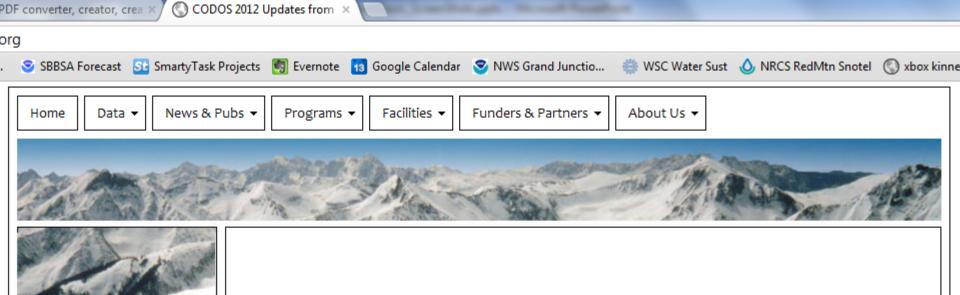
CODOS Program Funders	WY 2007	WY 2008	WY 2009	WY 2010	WY 2011	Actual WY 2012
Colorado River Water Conservation District	8,000	8,000	8,000	10,000	10,000	10,000
Southwestern Water Conservation District	5,000	5,000	4,000	5,000	5,000	5,000
Rio Grande Water Conservation District		3,000	4,000	5,000	5,000	5,000
Upper Gunnison River Water Conservancy Dist.		5,000	7,500	7,500	7,500	7,500
Northern Colorado Water Conservancy District			1,500	2,500	2,000	
Tri-County Water Conservancy District	1,000	1,000	1,500	2,500	2,500	2,500
Animas-La Plata Water Conservancy District			500	600	600	750
Dolores Water Conservancy District				600	600	750
Denver Water	2,500	2,500	2,500	5,000	5,000	5,000
Bureau of Reclamation – Western Colorado Area			5,000	7,500	7,500	
Bureau of Reclamation – Lower Colorado Region				7,500	10,000	10,000
Bureau of Reclamation – Eastern Colorado Area					2,500	
Bureau of Reclamation – Albuquerque Area						
Western Water Assessment – Univ of Colorado			20,072			
Colorado Water Conservation Board				28,034	15,000	25,000
City of Grand Junction					2,500	2,500
TOTAL	16,500	24,500	54,572	81,734	75,700	74,000



May 26, 2010 – Grizzly Peak Snotel, A-Basin







COLORADO DUST-ON-SNOW: WATER YEAR 2012 UPDATES

Welcome! This portion of our website is **exclusively for CODOS stakeholder organizations**. You won't find a link to this page from our public website. However, you are welcome to share the link to this page with your colleagues and partners: <u>dust.snowstudies.org</u>.

UPDATES & ALERTS BY DATE

- April 13, 2012: Summary of April 4-11 Circuit
- April 12, 2012: <u>McClure Pass</u>, <u>Wolf Creek Pass</u>, & <u>Spring Creek Pass</u> Updates
- April 10, 2012: <u>Grizzly Peak</u>, <u>Berthoud Summit</u>, Willow Creek Pass & Rabbit Ears Pass Updates
- April 9, 2012: <u>Park Cone</u> & <u>Hoosier Pass</u> Updates
- April 8, 2012: <u>Senator Beck Basin Update</u>
- April 6, 2012: <u>Dust alert [D8]</u>
- April 5, 2012: Grand Mesa update
- April 2, 2012: <u>Dust-on-dust Alert [D7]</u>
- March 28, 2012: <u>Update #3</u>
- March 26, 2012: Dust event underway [D6].

CODOS Quick Links

- About CODOS
- <u>Updates & Analysis</u> (prior years)
- <u>Dust Log & Wind Roses</u>
- <u>SNOTEL datasets</u>
- <u>Stakeholder Funding</u>
- Map of CODOS Sites
- Press Releases & Articles
- <u>Photo Gallery</u>



CODOS 2012 Updates > April 4-11 Update > Berthoud Summit

CODOS UPDATE FOR BERTHOUD SUMMIT: VISITED APRIL 10, 2012

Summary | Snowpack | Melt Rate | Forecast | Stream Flows | Earlier Updates

SUMMARY

Sustained periods of unseasonably warm air temperatures and exposed dust at the snowpack surface during late March and early April 2012 have, together, initiated accelerating rates of snowmelt and SWE loss at some, but not all, CODOS Snotel sites. Some CODOS Snotel sites report significant declines in SWE approaching the lowest values in the period of record (for a given date) or even falling outside



of the historic range. Those sites may have experienced Peak SWE for WY 2012 in early or mid-March. Recent CODOS snowpits near those CODOS sites mirror those losses of SWE.

In contrast, other CODOS Snotel sites and CODOS snowpits show only small losses of SWE. At those sites, energy inputs from warm air and direct absorption of solar energy by dust at the snowpack surface was consumed in warming the snowpack towards an isothermal state at 0° C, as a precursor to the loss of SWE and onset of snowmelt runoff. Since our prior site visit on March 14 the snowpack at the Berthoud Summit CODOS site has lost all cold content and is now isothermal.

The National Weather Service expects warming weather in the Colorado mountains through Wednesday with strong SW'ly winds developing on Wednesday afternoon ahead of a cooler but largely dry airmass. Unsettled and cooler weather will finish the week and run through the weekend, including chances for rain and/or snow showers each day.

😴 SBBSA Forecast 🔄 SmartyTask Projects 関 Evernote 🔢 Google Calendar 😒 NWS Grand Junctio...

🛞 WSC Water Sust 💧 NRCS RedMtn Snotel

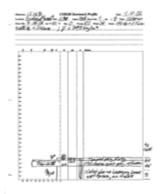
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SNOWPACK DISCUSSION

The snowcover at the Berthoud Summit CODOS site has undergone complete warming to 0°C thoughout since our March 15 snowpit and is beginning to more rapidly melt. We walked to this site on dry ground for much of the approach. Dust loading at the Berthoud Summit site is similar to that observed at Grizzly Peak - more intense than at Hoosier Pass, but still less intense than at our Senator Beck Basin study sites. Reduction in snow surface albedo has been sufficient to absorb additional solar energy at the snowpack surface and contribute to warming and ablation of the snowcover. As previously discussed, the Berthoud Summit Snotel site is in an open meadow, unshaded by the adjoining forest. As a consequence, Snotel snowmelt rate and snowpack depth data fully capture the influence of direct radiative forcing when snow albedo is lowered by dust, in contrast to other, shaded Snotel sites. Our CODOS snowpit site is located immediately in front of the Snotel station, in the same open meadow.

On March 15th the snowpack at our Berthoud Summit CODOS snowpit site was 44" (112 cm) deep and most of the snowpack consisted of very weak "depth hoar" grains; mean snowpack temperature was -2.6°C. Dust event D4 was clearly evident on the snow surface at the snowpit and in terrain around Berthoud Pass. SWE content in the snowpit was 13.2" (336 mm) and mean density of the snowpack was 308 kg/m³ (30.8% water content). As a result of the subsequent, prolonged period of warm, dry, and sunny weather, and some additional small reductions in snow albedo from additional dust events, the snowpack on April 10th was fully isothermal (0°C), with wet or very wet snow throughout. Total SWE in this pit was 8.7" (220 mm), a loss of 4.6" (116 mm) since March 15th; total snow depth was down to 20.5" (52 cm) and density had risen to 393 kg/m³ (39.3% water content).

April 10, 2012:



Pit profile

March 15, 2012:





Completed pit



Pit with Snotel in background





/dust/Berthoud/index.html



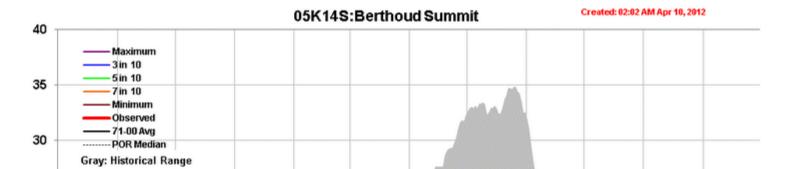
MELT RATE

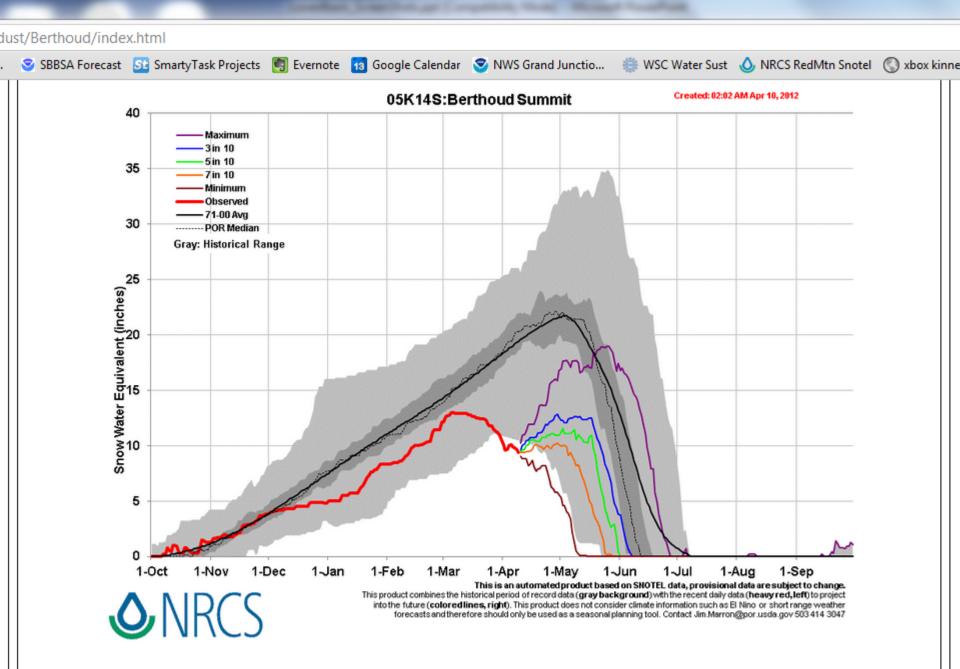
Berthoud Summit Snotel has reported a loss of 3.8" of SWE since our last site visit on March 15, 2012, not long after what may have been Peak SWE for WY 2012. This melt rate falls well short of the mean daily loss of SWE at Berthoud Summit observed in prior snowmelt seasons, as shown in the table below. However, even at current melt rates, SWE values may remain below the lowest quartile for the duration of WY 2012.

Berthoud Summit SNOTEL Snowmelt Season Summary Data

						Adjusted	
					Post-Peak	Daily	Period
		Date	Peak	Days	Added	Mean Loss	Mean
		Peak SWE	SWE	to SAG	SWE	SWE	Temp
WY 2006		4/21/2006	24.0	41	3.8	0.68	3.5
WY 2007		4/27/2007	22.2	46	4.5	0.58	4.4
WY 2008		5/16/2008	24.4	34	1.4	0.76	5.8
WY 2009		4/20/2009	24.7	50	5.2	0.60	4.0
WY 2010		5/16/2010	24.5	23	0.6	1.09	6.6
WY 2011		5/26/2011	34.8	35	2.0	1.05	8.4
	Max	5/26	34.8	50	5.2	1.09	8.4
	Min	4/20	22.2	23	0.6	0.58	3.5
	Range	37	12.6	27	4.6	0.51	4.9

Adjusted Daily Mean Loss SWE rates include additional SWE received after date of Peak SWE





STREAM FLOWS

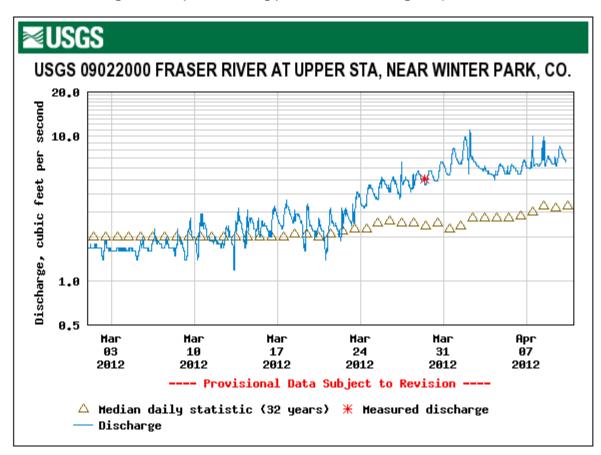
Streamflow behavior at the USGS Fraser River Upper Staion near Winter Park gauge reports a brief decline in discharge in early April, after a significant

ONRCS

This is an automated product based on SNOTEL data, provisional data are subject to change. This product combines the historical period of record data (gray background) with the recent daily data (heavy red, left) to project into the future (colored lines, right). This product does not consider climate information such as El Nino or short range weather forecasts and therefore should only be used as a seasonal planning tool. Contact Jim Marron@por.usda.gov 503 414 3047

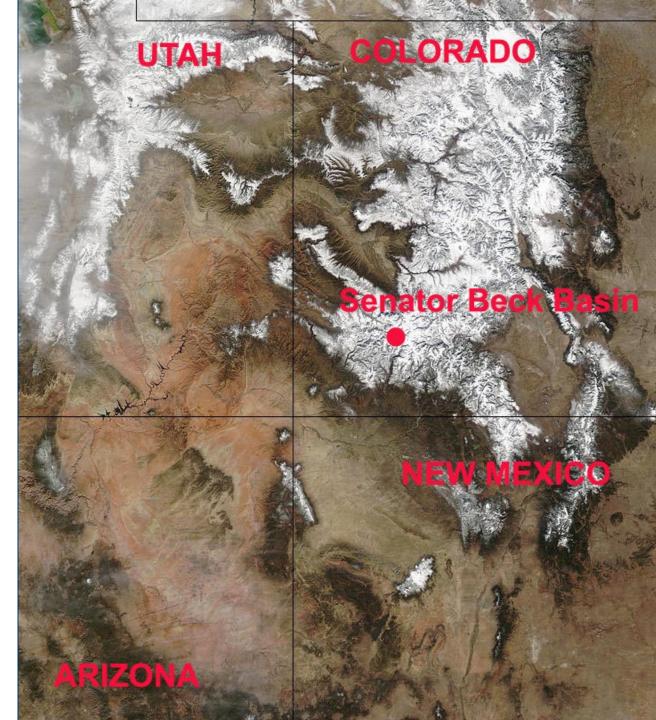
STREAM FLOWS

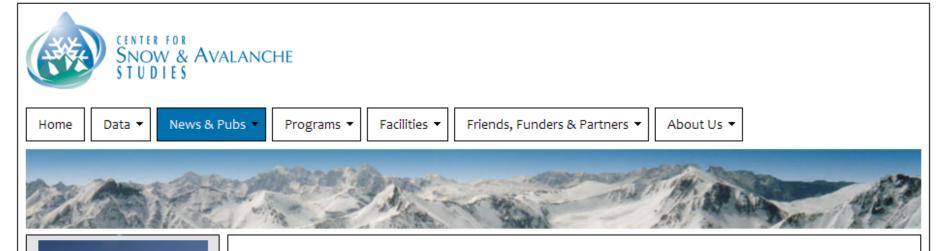
Streamflow behavior at the USGS Fraser River Upper Staion near Winter Park gauge reports a brief decline in discharge in early April, after a significant surge in late March. Flows in late March were high, compared to median levels at that gauge, for that period. Unsettled weather in early April ushered in cooler air and cloudier skies for the following several days. Interestingly, flows stabilized during that period at more than double the median values.



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Senator Beck Basin research has broad application









Senator Beck Study Plot (alpine)

CSAS-ASSISTED SCHOLARLY PUBLICATIONS

Naud, C. M., J. R. Miller, and C. Landry (2012), Using satellites to investigate the sensitivity of longwave downward radiation to water vapor at high elevations, J. Geophys. Res., 117, D05101, doi:10.1029/2011JD016917.

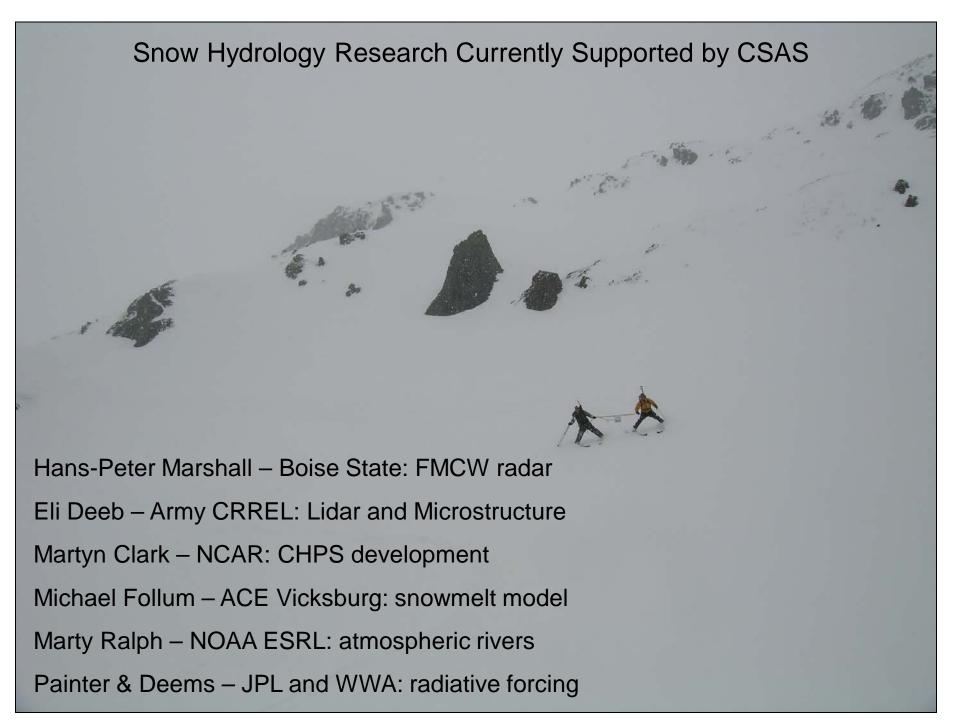
Marshall, H.P., C. Pielmeier, S. Havens, and F. Techel (2010), Slope-scale Snowpack Stability Derived from Multiple Snowmicropen Measurements and High-resolution Terrestrial FMCW Radar Surveys. Proceedings of the 2010 International Snow Science Workshop, Squaw Valley, California.

Simonson, S.E., E. Greene, S. Fasnacht, T. Stohlgren and C. Landry (2010) Practical Methods for Using Vegetation Patterns to Estimate Avalanche Frequency Magnitude. Proceedings of the 2010 International Snow Science Workshop, Squaw Valley, California.

Painter, T. H., J. Deems, J. Belnap, A. Hamlet, C. C. Landry, and B. Udall (2010), Response of Colorado River runoff to dust radiative forcing in snow, Proceedings of the National Academy of Sciences, published ahead of print September 20, 2010, doi:10.1073/pnas.0913139107.

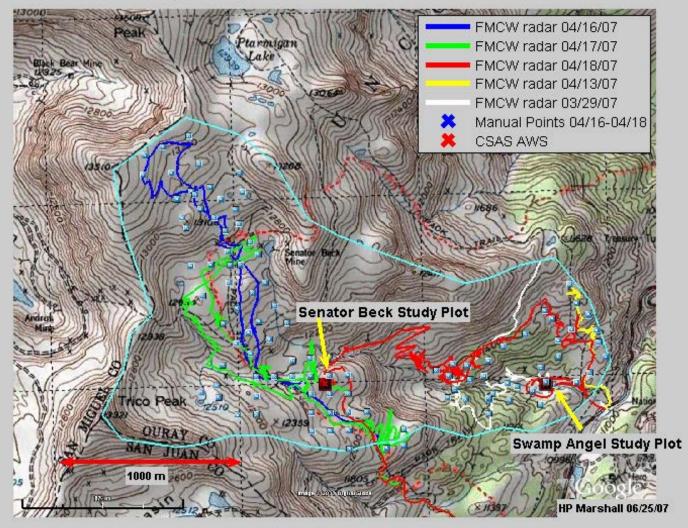
Lawrence, C. R., T. H. Painter, C. C. Landry, and J. C. Neff (2010), Contemporary geochemical composition and flux of aeolian dust to the San Juan Mountains, Colorado, United States, Journal of Geophysical Research, 115, G03007, doi:10.1029/2009JG001077.

Steltzer, H., C. Landry, T. H. Painter, J. Anderson, and E. Ayres. 2009. Biological consequences of earlier snowmelt from desert dust deposition in alpine landscapes. Proceedings of the National Academy of Sciences. 106: 11629-11634, doi_10.1073_pnas.0900758106.

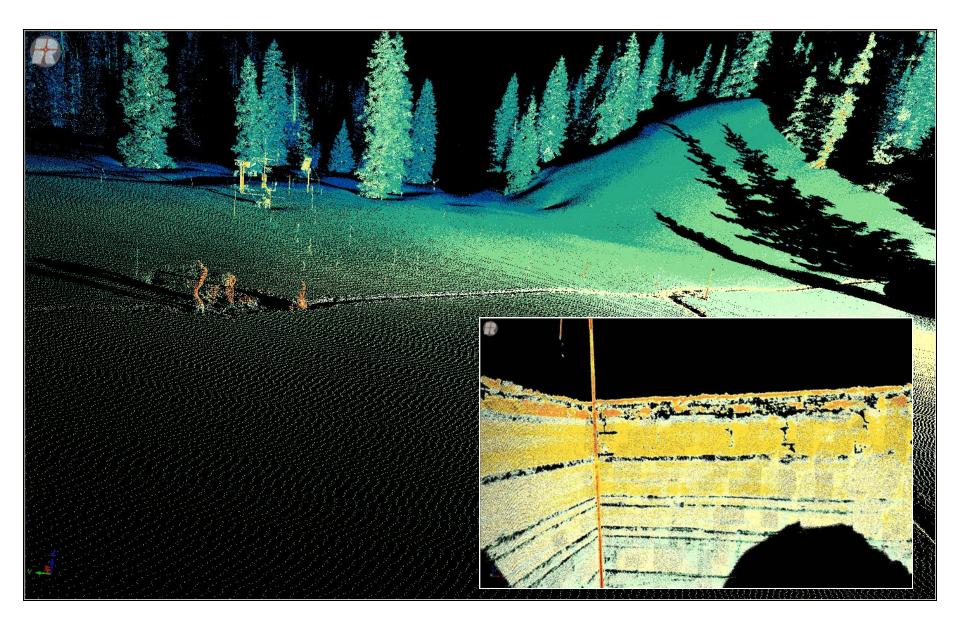


Hans Peter Marshall - FMCW Radar Development

Senator Beck Study Area, Center for Snow and Avalanche Studies



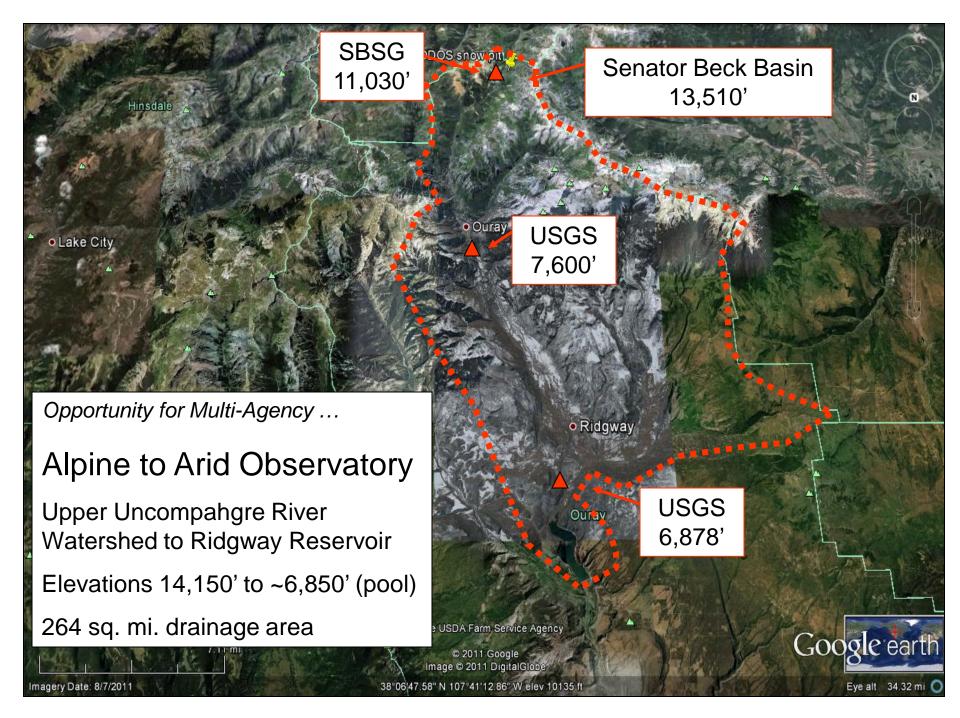
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Climate Change in Mountains

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Journal Details Home AGU Journals	JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 117, D05101, 12 PP., 2012 doi:10.1029/2011JD016917 Using satellites to investigate the sensitivity of longwave downward radiation to water vapor at high elevations	 E-Alert Sign-Up RSS Feeds Cread By Reference Tools Contact AGU
Article Resources		
	Key Points • PWV vs. q universal for large q but elevation dependent for low q	Keywords • high elevation
Full Text (HTML)	Key Points	Keywords • high elevation • longwave downward flux
Full Text (HTML) Full Text (PDF)	 Key Points PWV vs. q universal for large q but elevation dependent for low q 	Keywords high elevation
Full Text (HTML) Full Text (PDF) Purchase Article E-mail Abstract	Key Points • PWV vs. q universal for large q but elevation dependent for low q • MODIS PWV and CERES LDR as accurate at high as at low elevations	Keywords high elevation longwave downward flux precipitable water vapor
Full Text (HTML) Full Text (PDF) Purchase Article	Key Points • PWV vs. q universal for large q but elevation dependent for low q • MODIS PWV and CERES LDR as accurate at high as at low elevations • Satellites observe high sensitivity of LDR to changes in q in dry locations Catherine M. Naud	Keywords • high elevation • longwave downward flux • precipitable water vapor • satellite observations • specific humidity

However, in situ observations in these often remote locations are too sparse to determine the feedbacks responsible for enhanced warming rates. One of these feedbacks is associated with the sensitivity of longwave downward radiation (LDR) to changes in water vapor, with the sensitivity being particularly large in many high-elevation regions where the average water vapor is often low. We show that satellite retrievals from the Moderate Resolution Imaging Spectroradiometer (MODIS) and Clouds and the Earth's Radiant Energy System (CERES) can be used to expand the current ground-based observational database and that the monthly averaged clear-sky satellite estimates of humidity and LDR are in good agreement with the well-instrumented Center for Snow and Avalanche Studies ground-based site in the southwestern Colorado Rocky Mountains. The relationship between MODIS-retrieved precipitable water vapor and surface specific humidity across the contiguous United States was found to be similar to that previously found for the Alps. More important, we show that satellites capture the nonlinear relationship between LDR and water vapor and confirm that LDR is especially sensitive to changes in







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